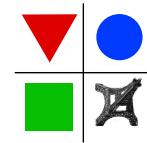


# A multilayer model for non-hydrostatic multiscale free-surface flows

Stéphane Popinet

Institut *d'Alembert*

CNRS / Sorbonne Université



***d'Alembert***  
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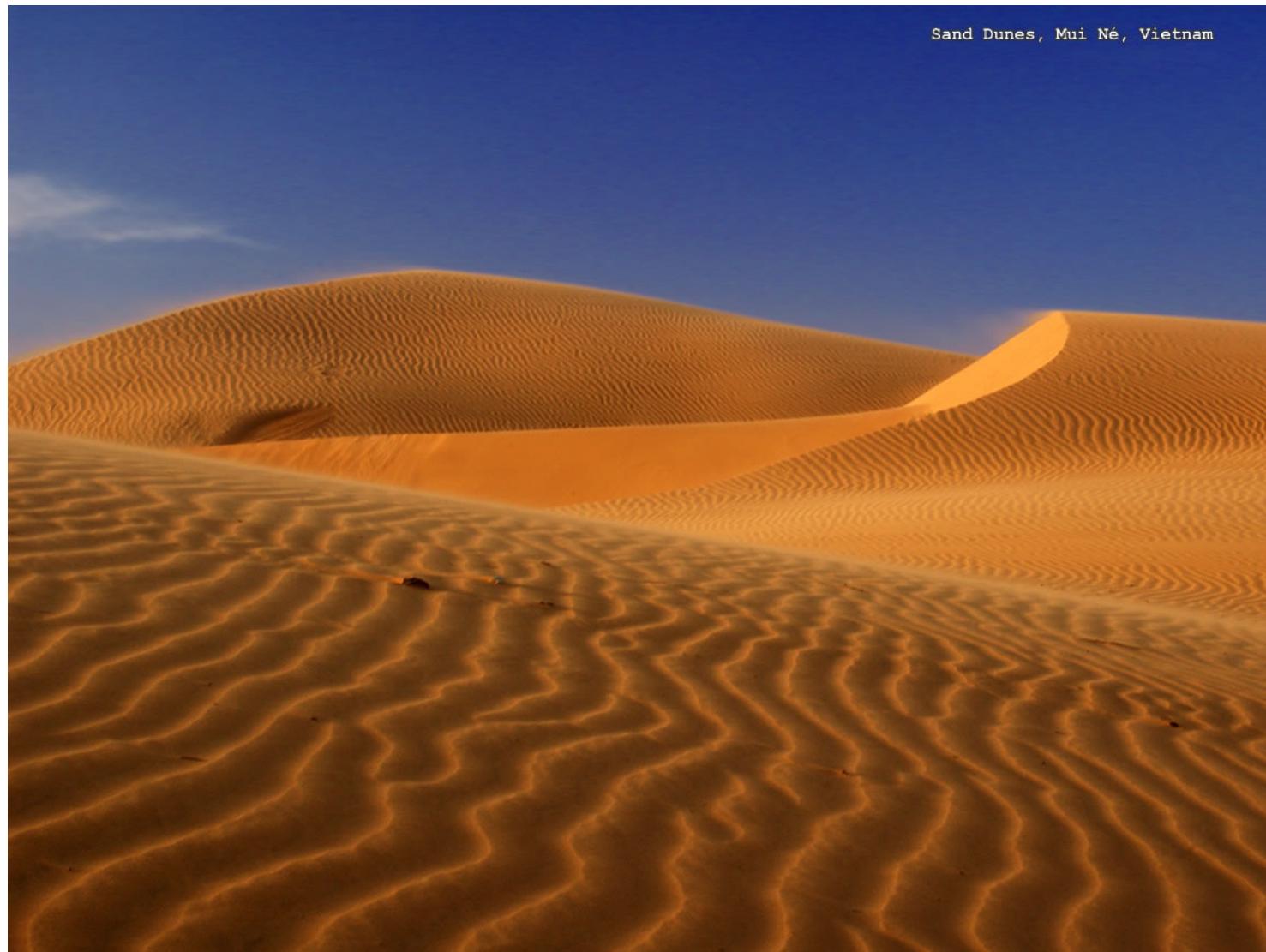


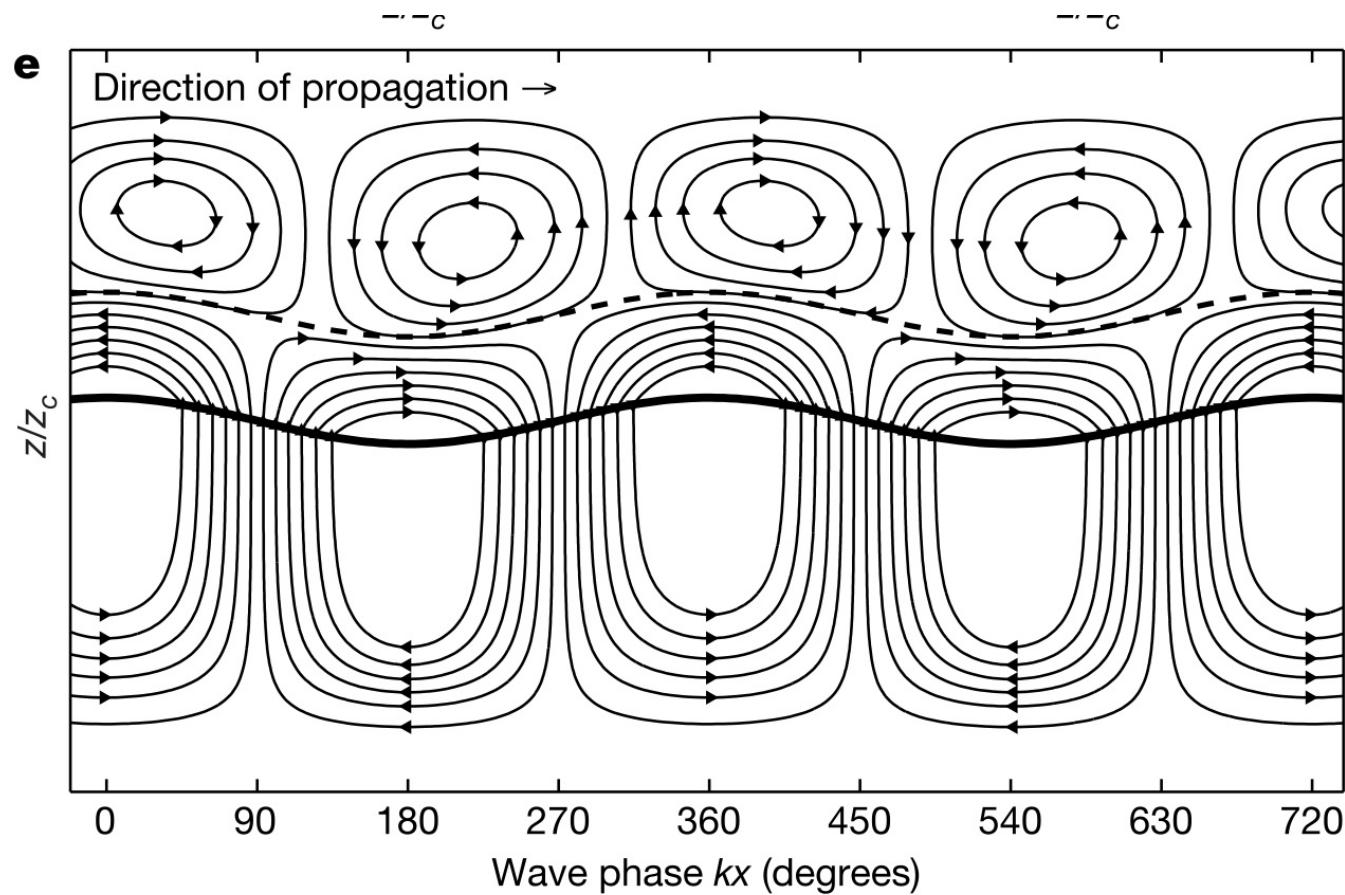


# Avalanches / granular materials

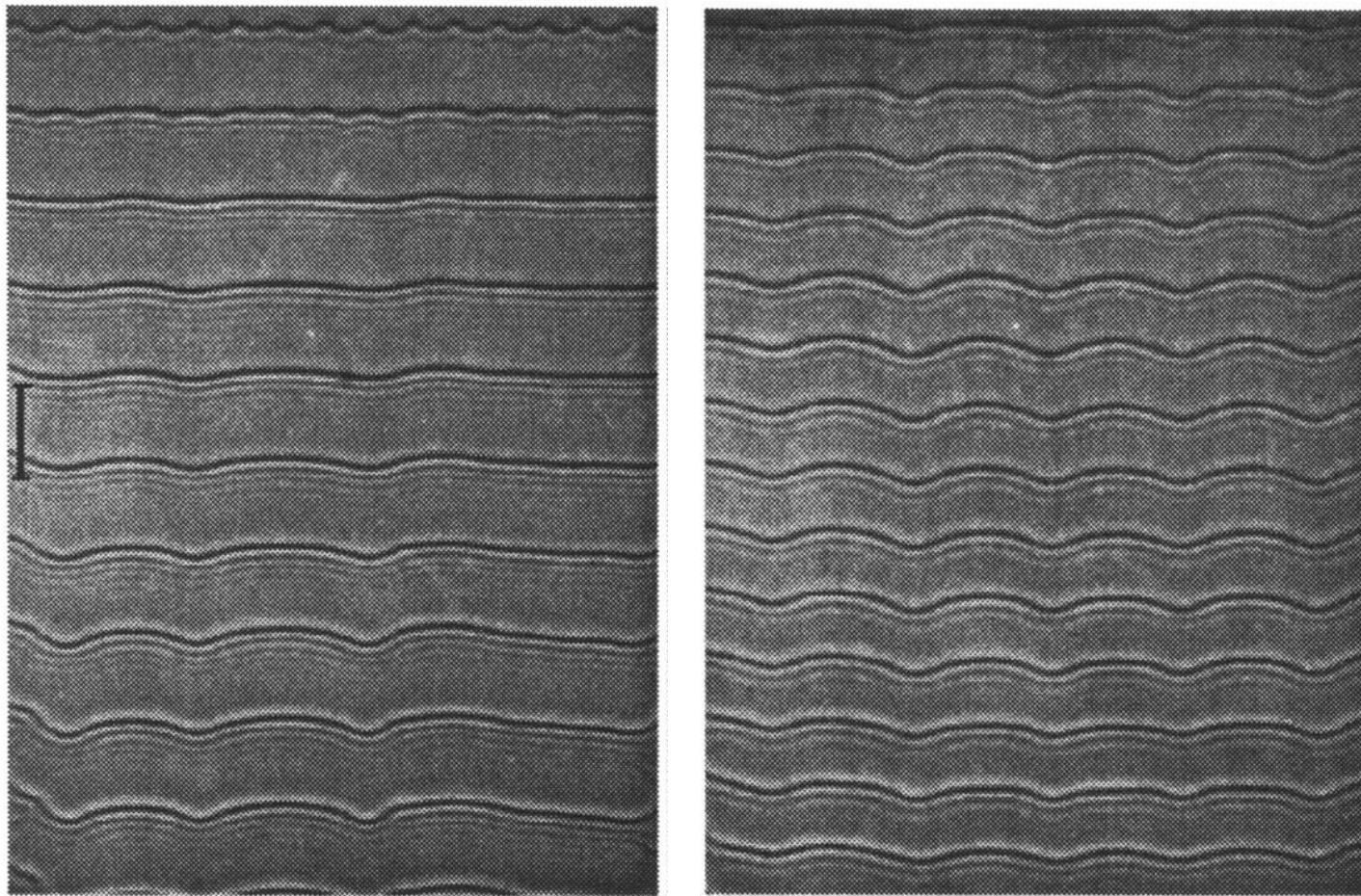
4/26

Sand Dunes, Mui Né, Vietnam





Hristov et al, *Dynamical coupling of wind and ocean waves through wave-induced air flow*, Nature, 2003.



Piotr Leonidovich Kapitza, "Wave flow of thin layers of a viscous fluid," in Collected papers of P.L. Kapitza, D. Ter Haar, Ed., pp. 662–689. Pergamon, 1948.

$$\begin{aligned}\partial_t h + \nabla \cdot (h \mathbf{u}) &= 0, \\ \partial_t (h \mathbf{u}) + \nabla \cdot (h \mathbf{u} \mathbf{u}) &= -g h \nabla \eta, \\ \eta &= z_b + h\end{aligned}$$

Adhémar Jean-Claude Barré de Saint-Venant, 1871.

- + Simple
- + System of conservation laws
- + Fully non-linear
- + Numerically efficient
- Vertical structure is entirely modelled
- Hydrostatic, no-vertical acceleration/momentum

Non-dispersive waves:  $c = \sqrt{g h}$



1797–1886

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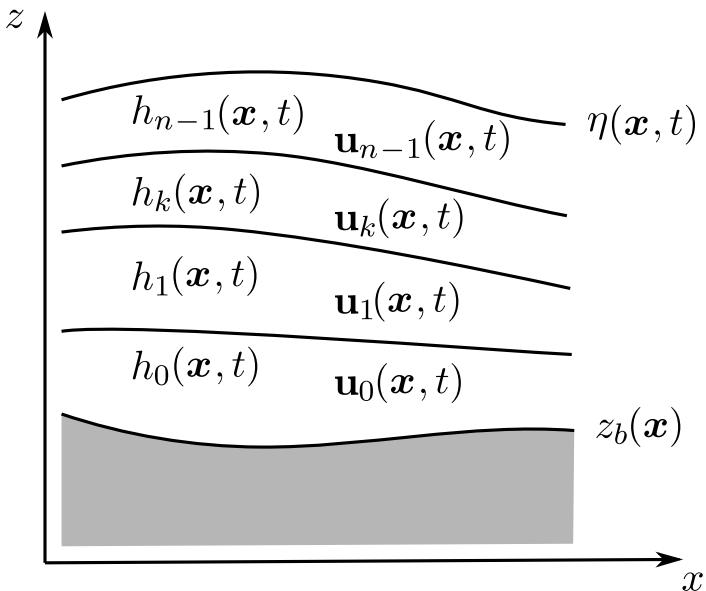


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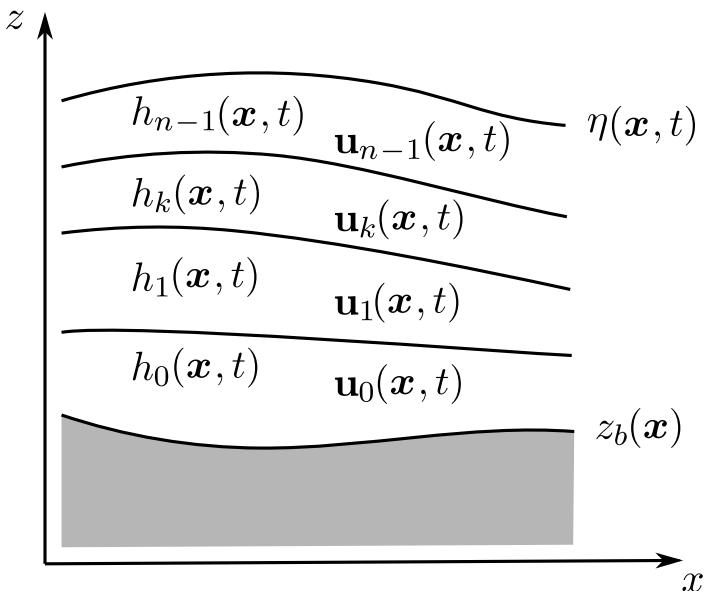


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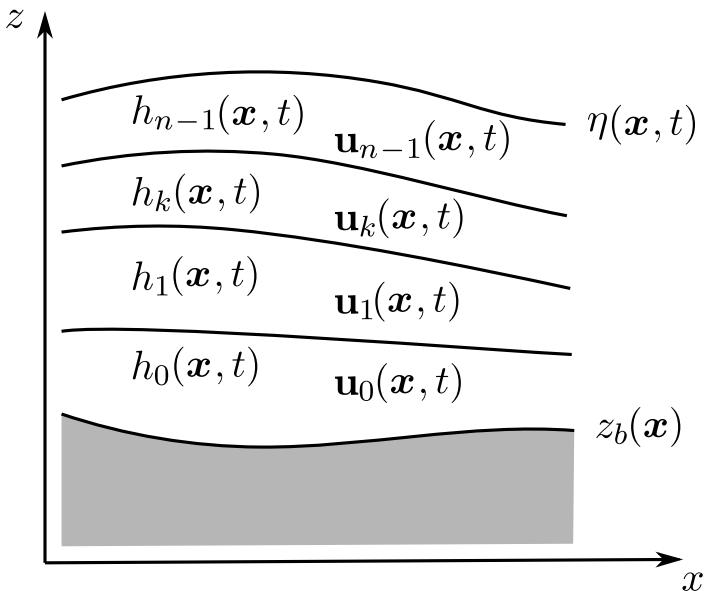


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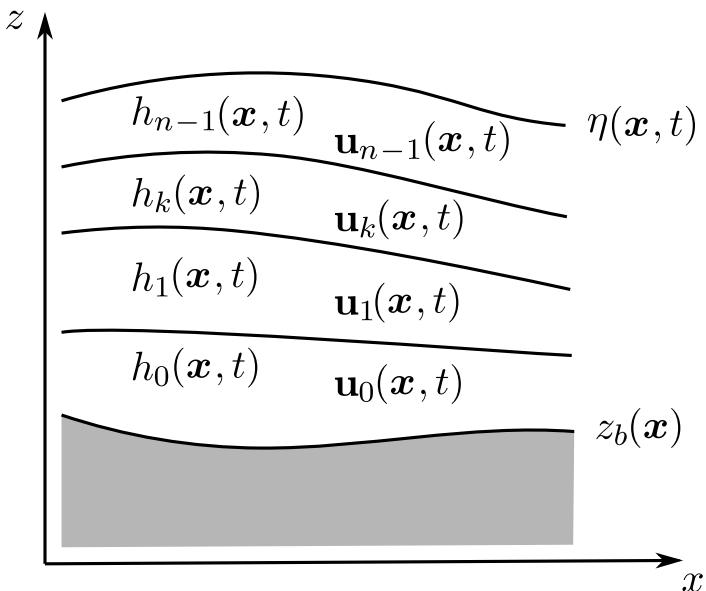


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M. BOUSSINESQ.

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Leonhard Euler, 1757 (without the free surface).

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- Link with Saint-Venant?
- How to represent  $\chi$ ? (VOF, Levelset, Lagrangian etc.)
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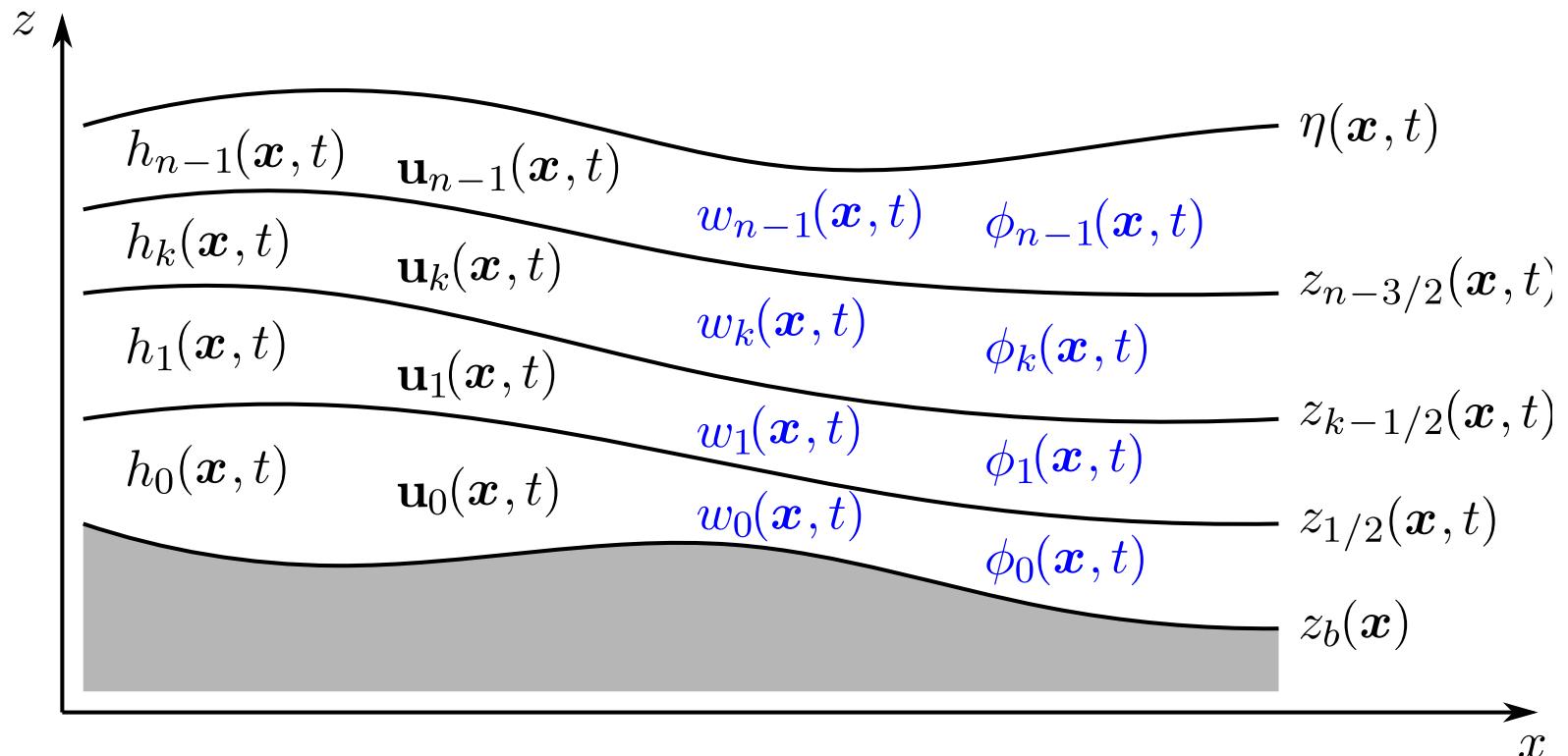
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- + Clearly connected with the simpler hydrostatic system
- + System of (discrete) conservation laws
- + Numerically efficient
- +– Vertically Lagrangian
- Too many unknowns!

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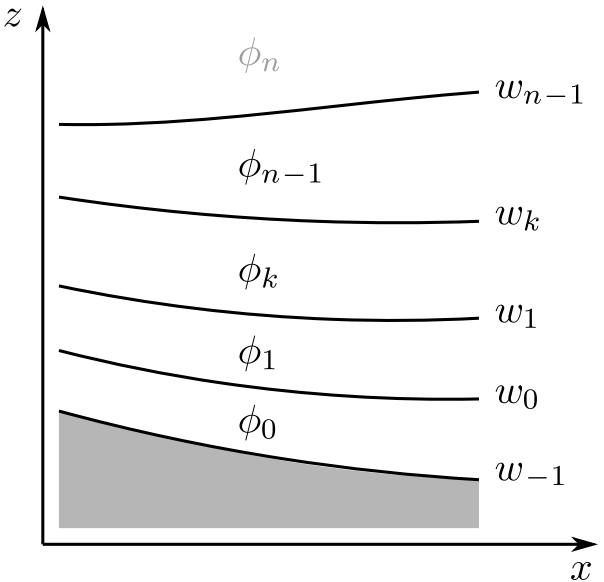
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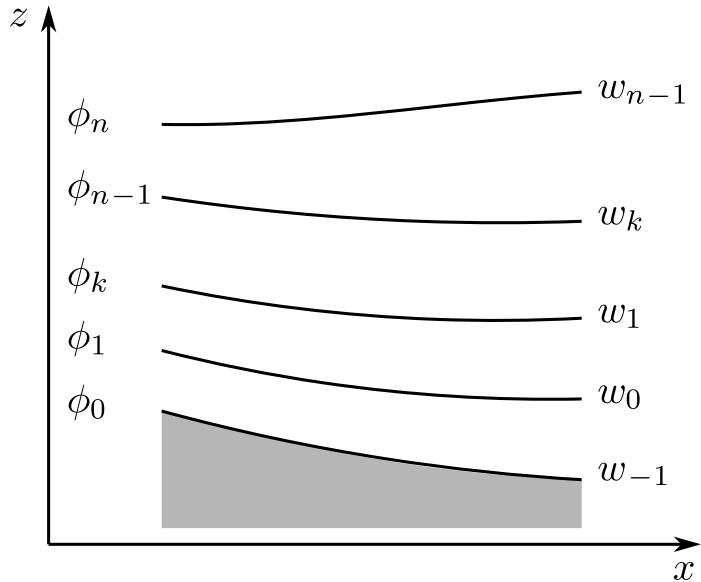
## Examples of possible vertical discretisations



1917–2008



Lorenz grid



Keller grid (box scheme)

Edward Norton Lorenz, *Energy and numerical weather prediction*, 1960.

H. B. Keller, *A new difference scheme for parabolic problems*. 1971.

Zijlema & Stelling, IJNMF, 2005.

Which one is best? → choose the best dispersion relation based on the *linearised perturbation system* ( $h_k = \bar{h}_k + h'_k e^{i(\hat{k}x - \omega t)}$ ,  $u_k = u'_k e^{i(\hat{k}x - \omega t)}$ , etc.)

$$-\omega h'_k + \bar{h}_k u'_k \hat{k} = 0,$$

$$-\bar{h}_k u'_k \omega = -g \hat{k} \bar{h}_k \sum h'_k - \bar{h}_k \hat{k} \frac{\phi'_{k+1} + \phi'_k}{2},$$

$$-\bar{h}_k \frac{w'_k + w'_{k-1}}{2} \omega i = \phi'_k - \phi'_{k+1},$$

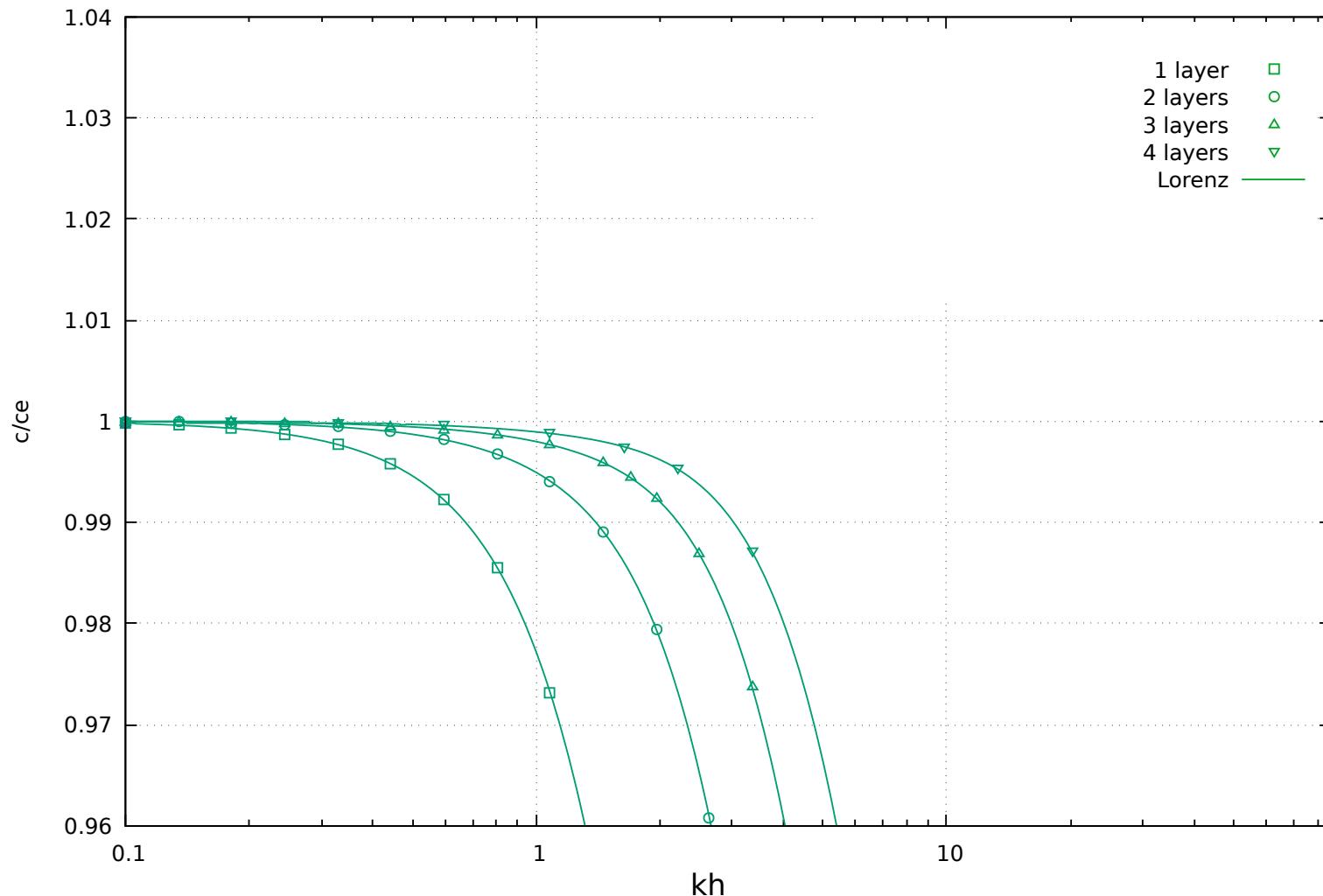
$$\bar{h}_k u'_k \hat{k} i + w'_k - w'_{k-1} = 0,$$

Computing the determinant then gives (for two layers)

$$\omega_2^2(\hat{k}) = 8 g \frac{h_0 h_1^2 \hat{k}^4 + (h_1 + h_0) \hat{k}^2}{3 h_0^2 h_1^2 \hat{k}^4 + 3 (3 h_0^2 + h_0 h_1 + h_1^2) \hat{k}^2 + 8}$$

# Dispersion relations

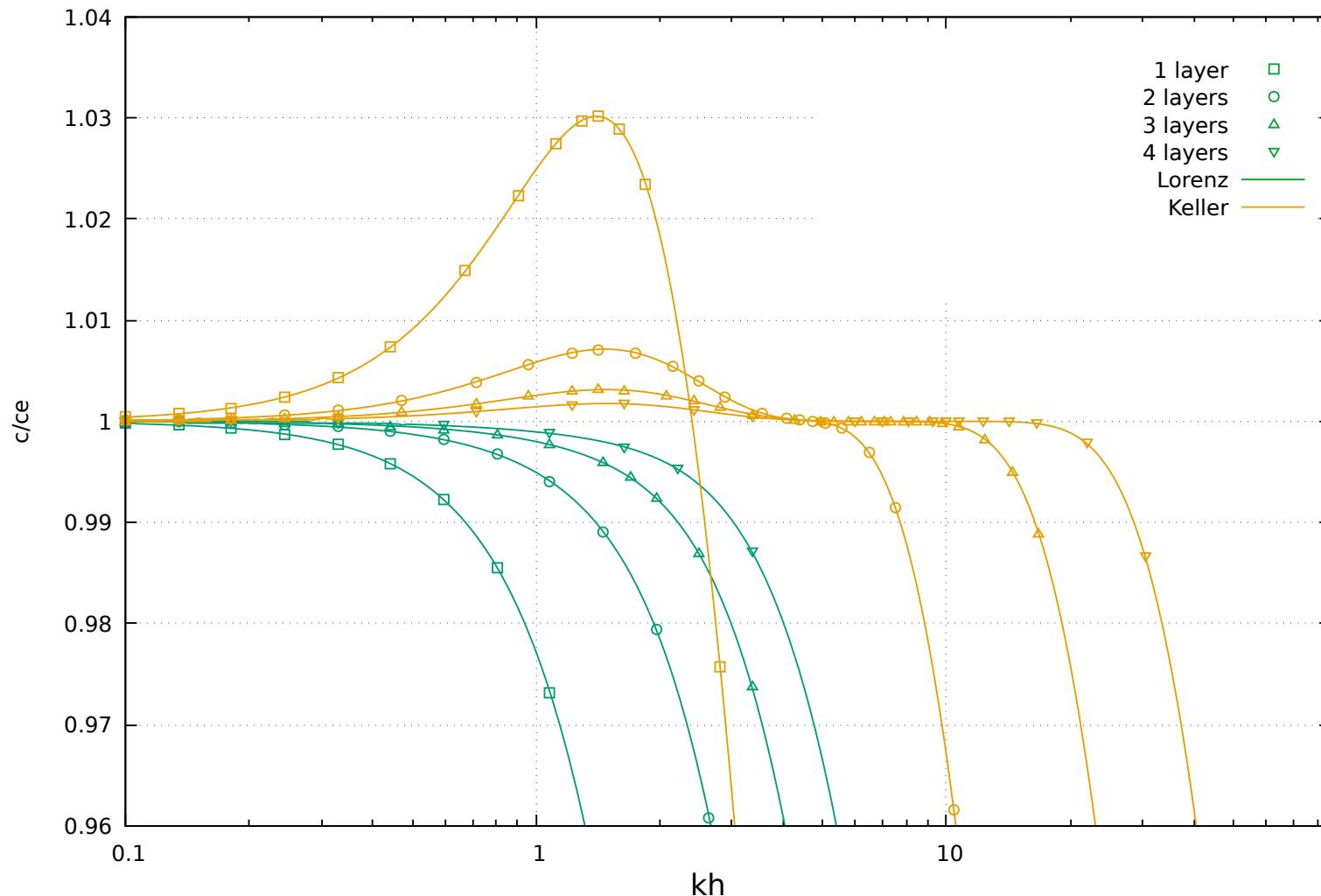
15/26



$$c_e^2 = \frac{g}{k} \tanh(k h)$$

# Dispersion relations

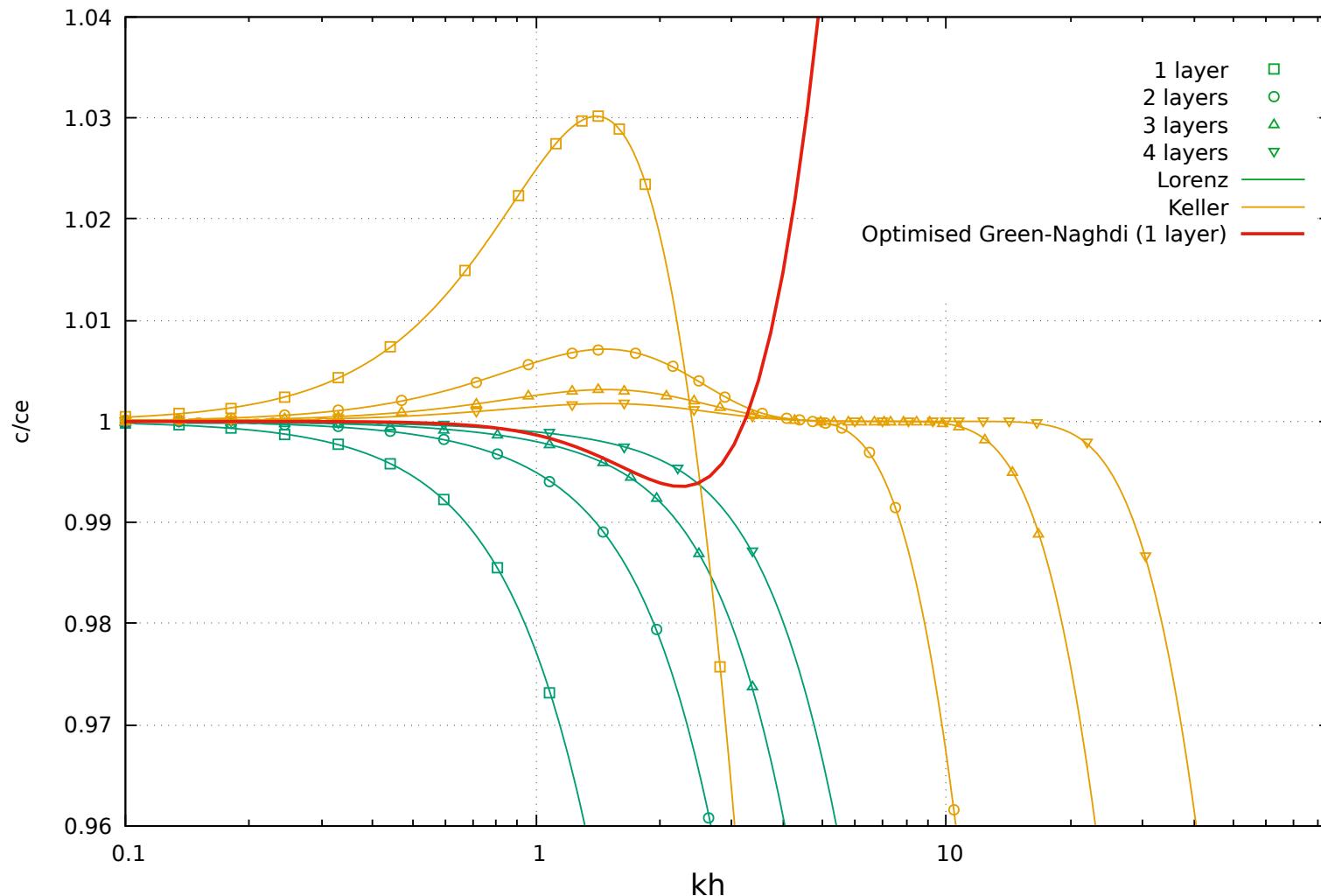
15/26



$$c_e^2 = \frac{g}{k} \tanh(k h)$$

# Dispersion relations

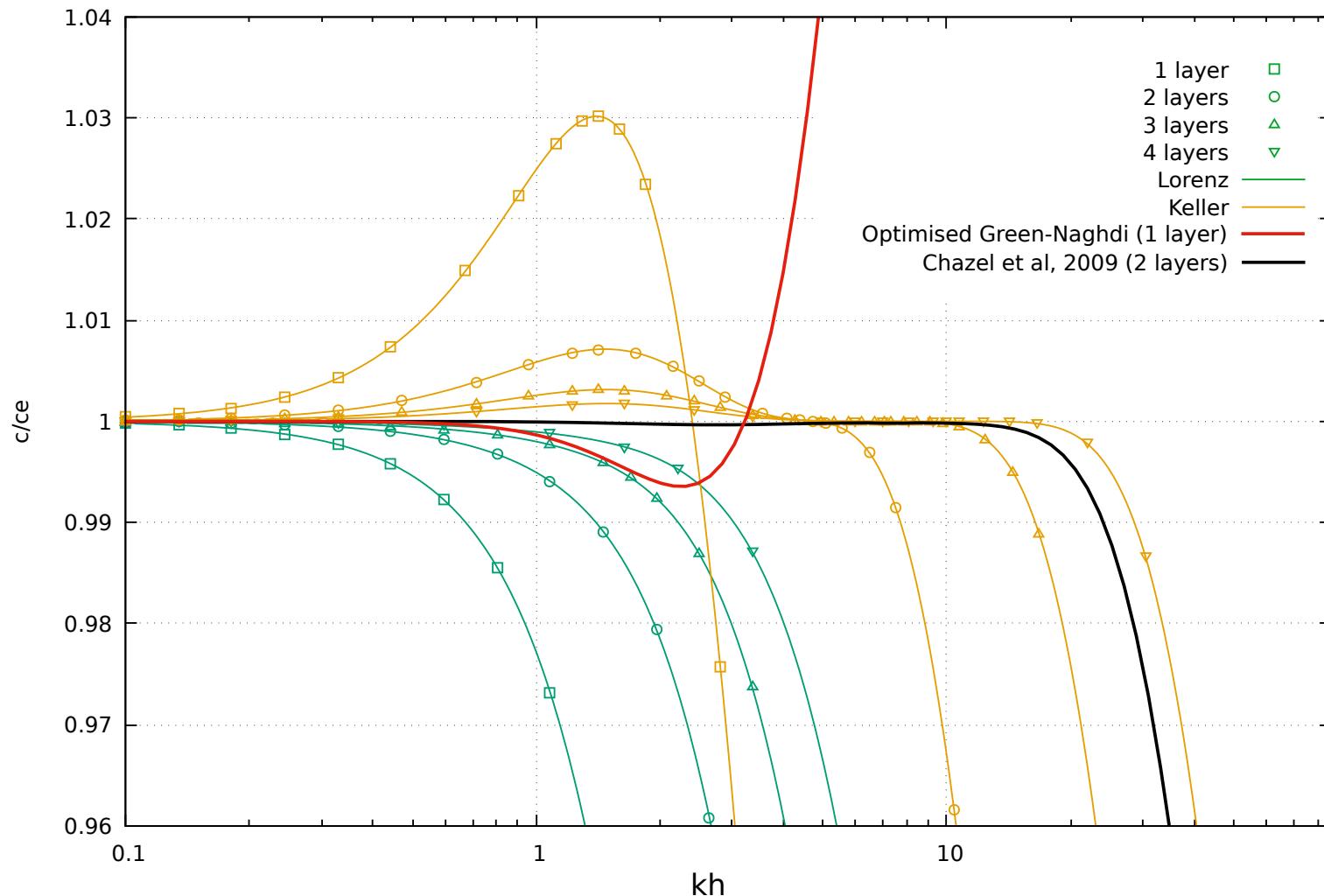
15/26



$$c_e^2 = \frac{g}{k} \tanh(k h)$$

# Dispersion relations

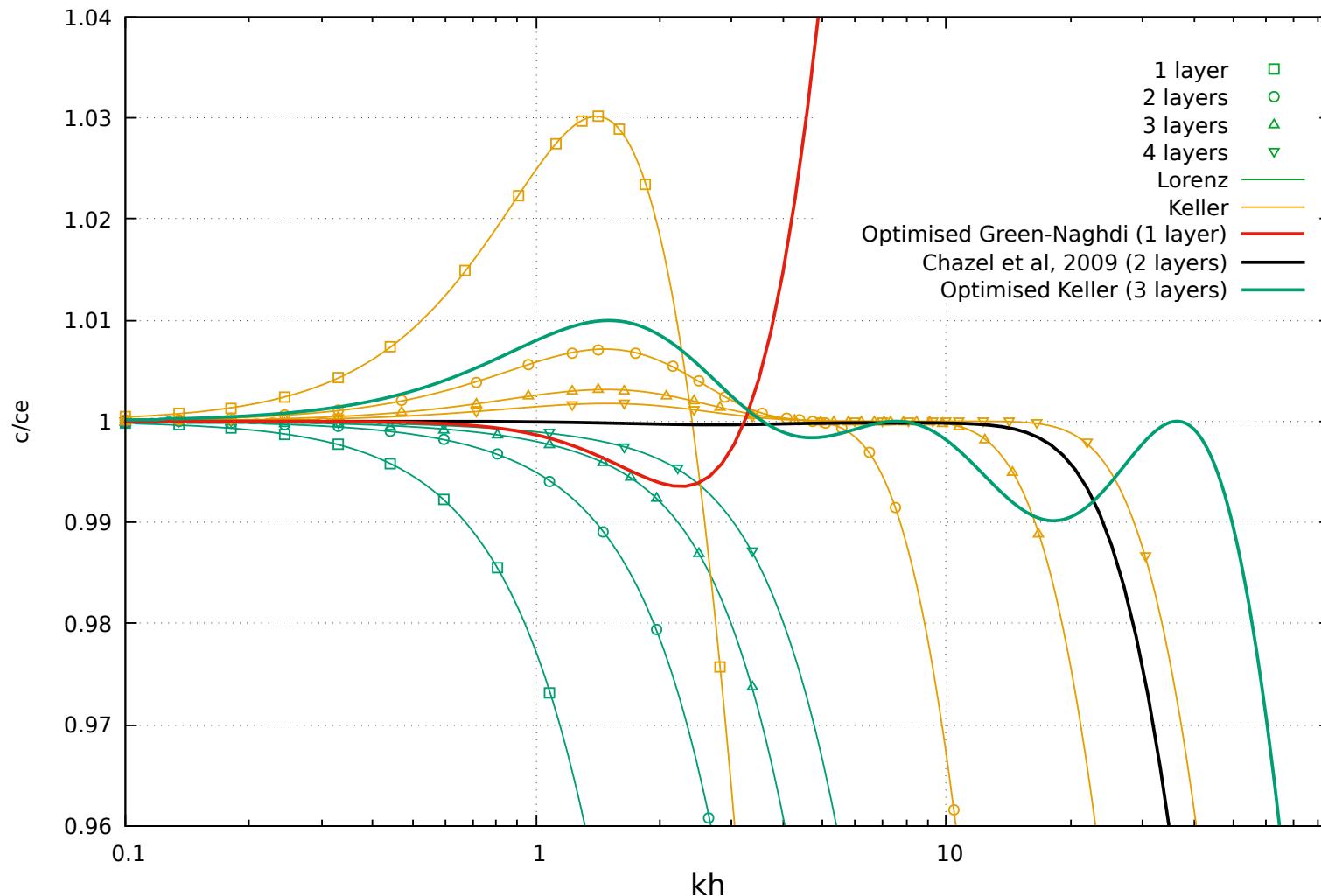
15/26



$$c_e^2 = \frac{g}{k} \tanh(k h)$$

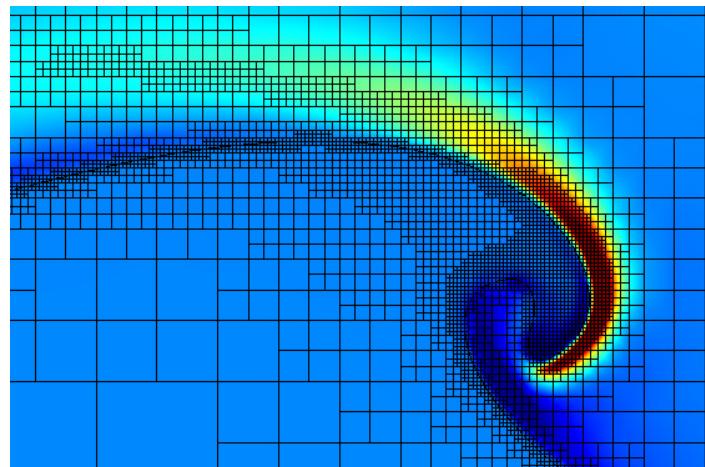
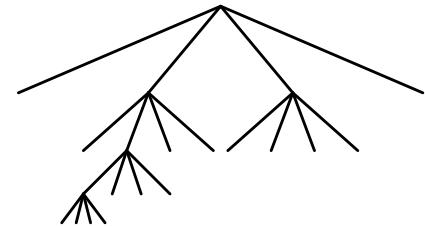
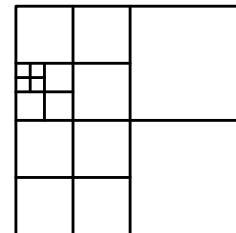
# Dispersion relations

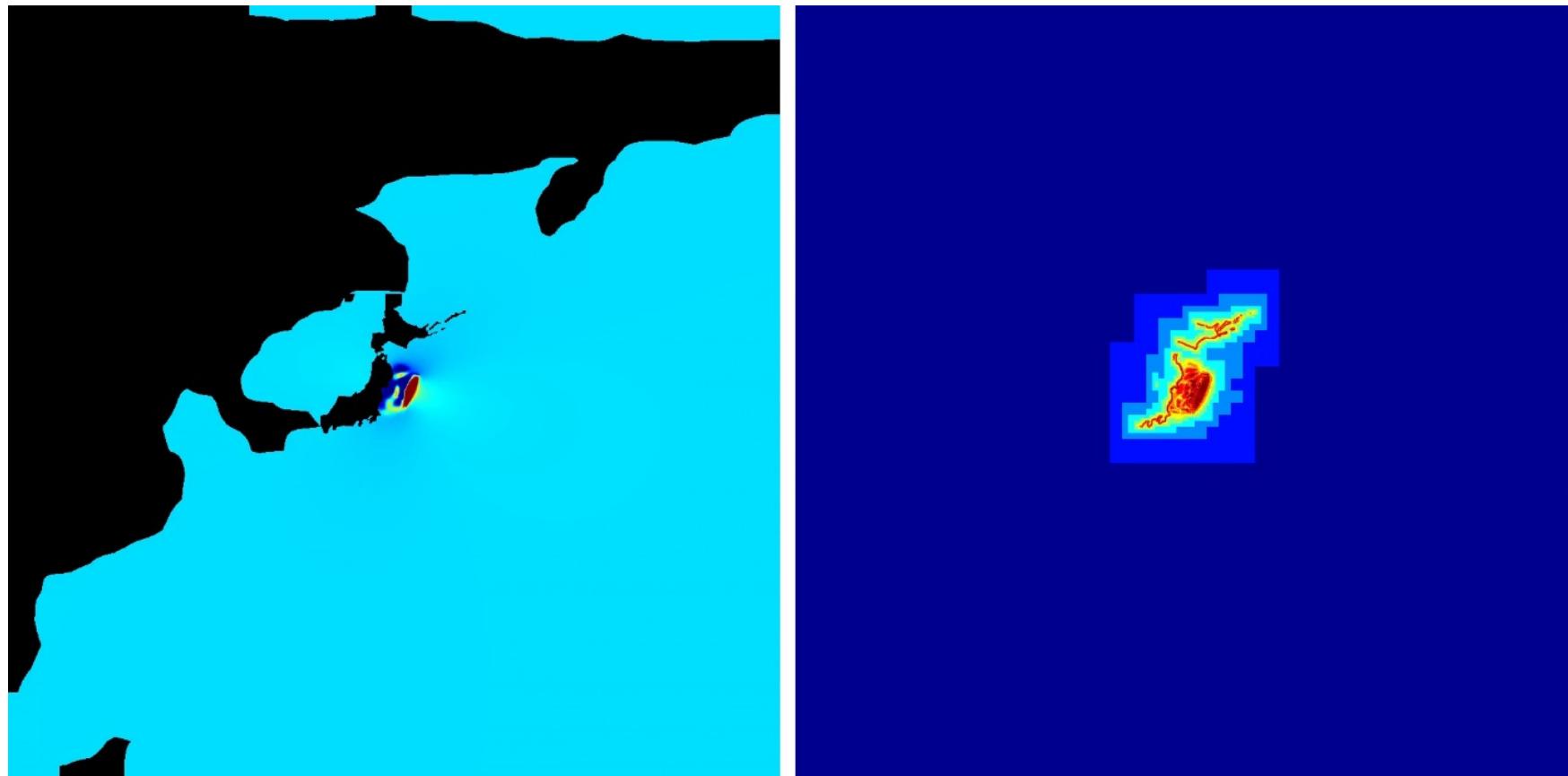
15/26



$$c_e^2 = \frac{g}{k} \tanh(k h)$$

- Basilisk: open-source, [basilisk.fr](http://basilisk.fr)
- Quad/octree adaptive mesh refinement
- Multigrid solver with (vertical) line relaxations for elliptic problems
- MPI/Open MP parallelism
- Load-balancing based on space-filling curves (Z-ordering) (scales to 1000s of cores)
- Decoupled (high-level) numerical scheme / (low-level) memory and loop traversal implementation
- Documented, reproducible and editable examples and test cases: “wikipedia of fluid mechanics”

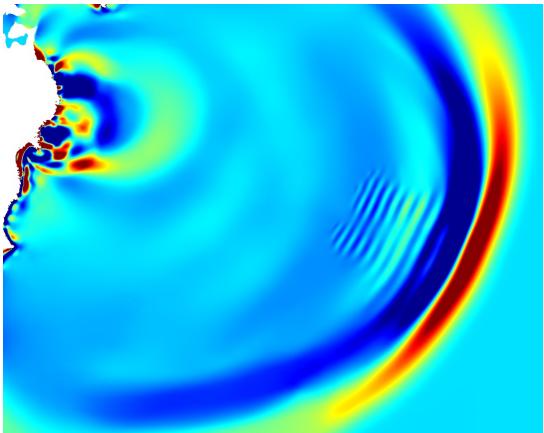




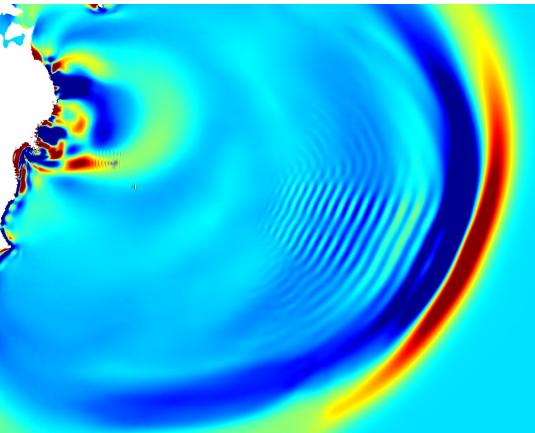
6000×8000 km, spatial resolution: 1–250 km

See also Popinet, 2015 (Green-Naghdi) and Popinet, 2012 (Saint-Venant)

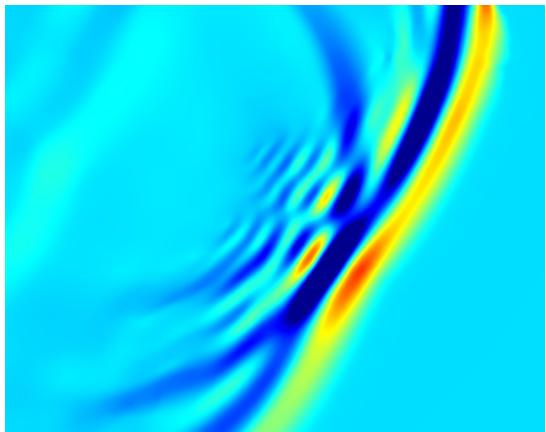
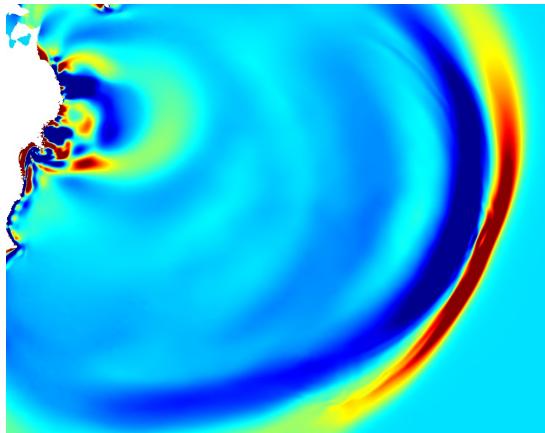
Serre–Green–Naghdi



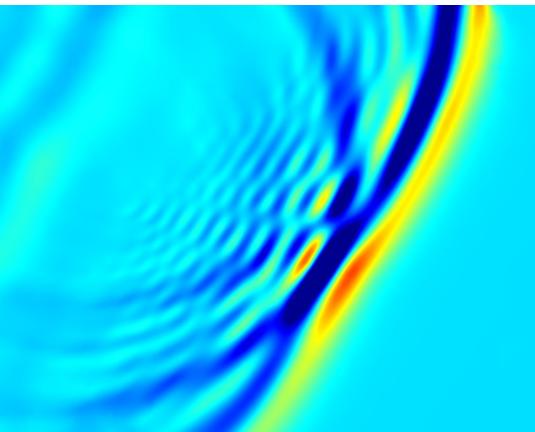
One-layer non-hydrostatic



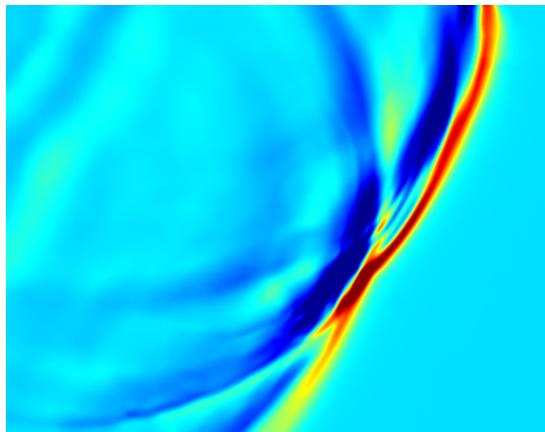
One-layer hydrostatic



CPU runtime: 7h30



CPU runtime: 4h15

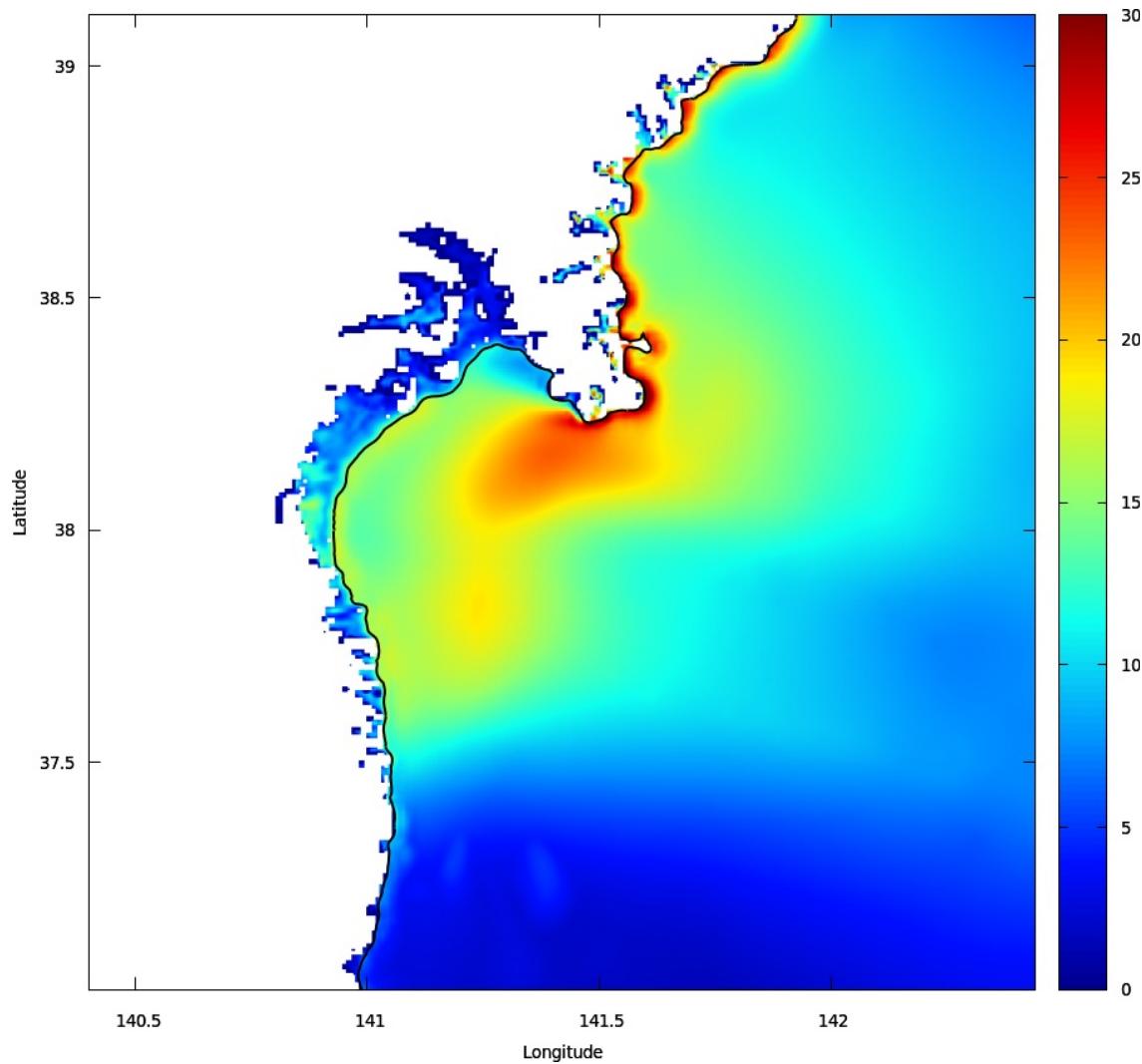


CPU runtime: 5h15

Computational speed  $\approx 250\ 000$  points  $\times$  timesteps / sec

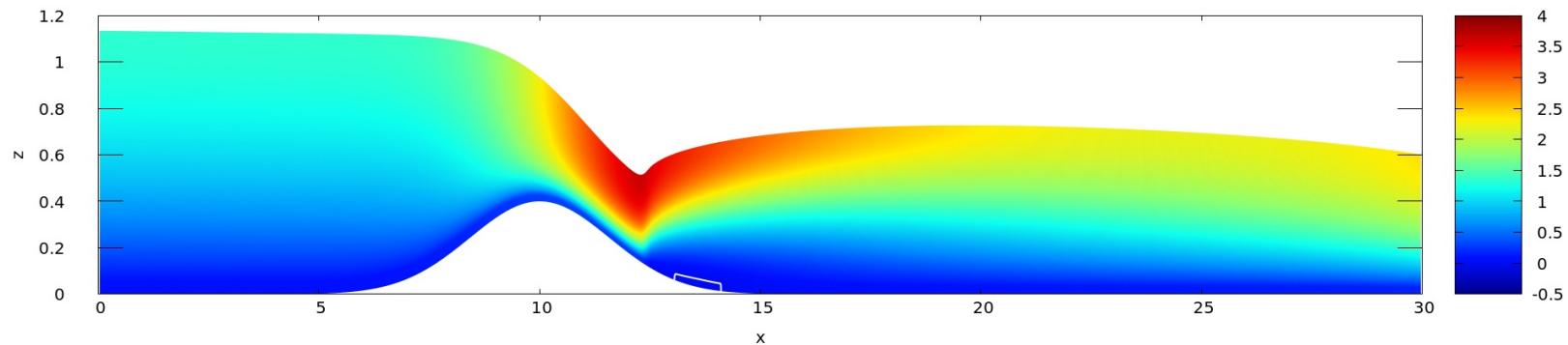
# Detail of flooding

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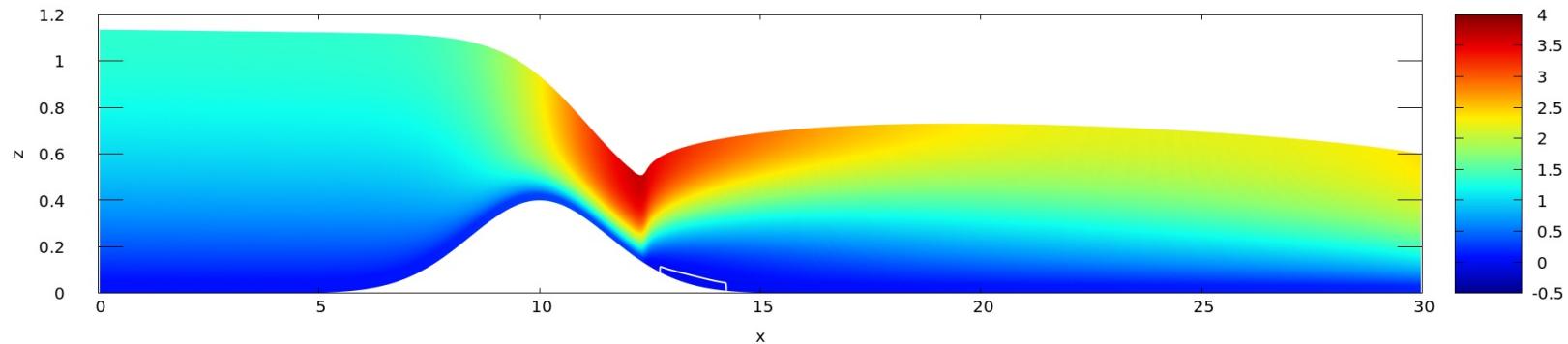


Sendai plain:  $140 \times 200$  km

## Horizontal velocity

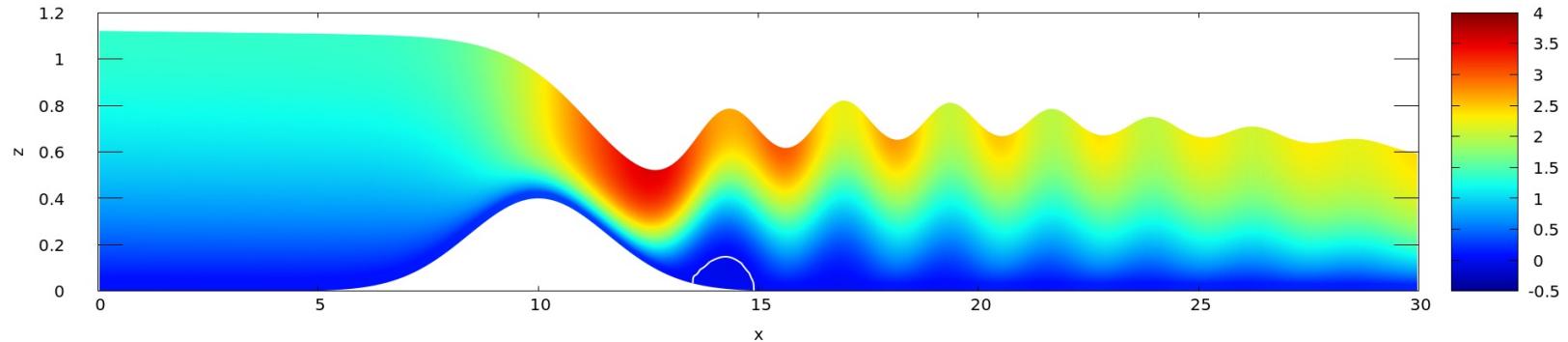


## Multilayer hydrostatic

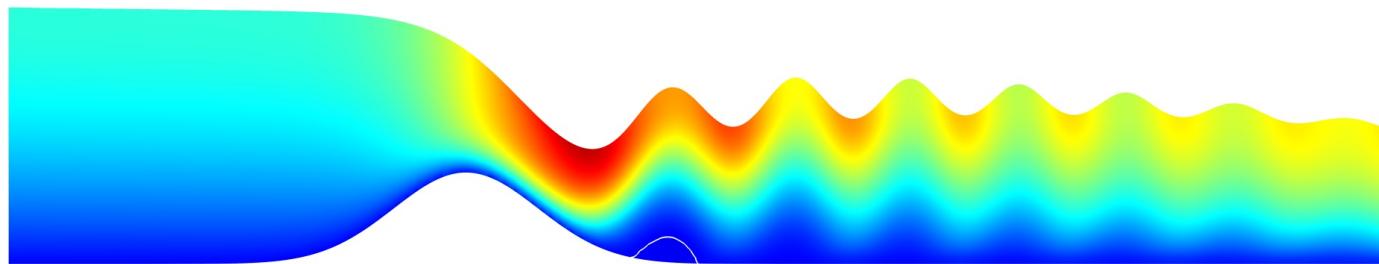


## Riemann solver

Horizontal velocity

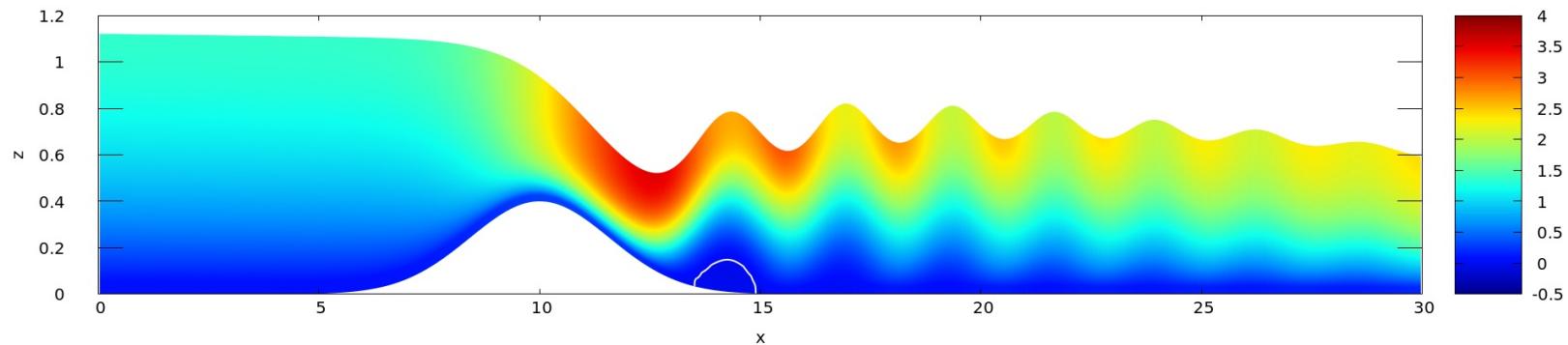


Multilayer non-hydrostatic

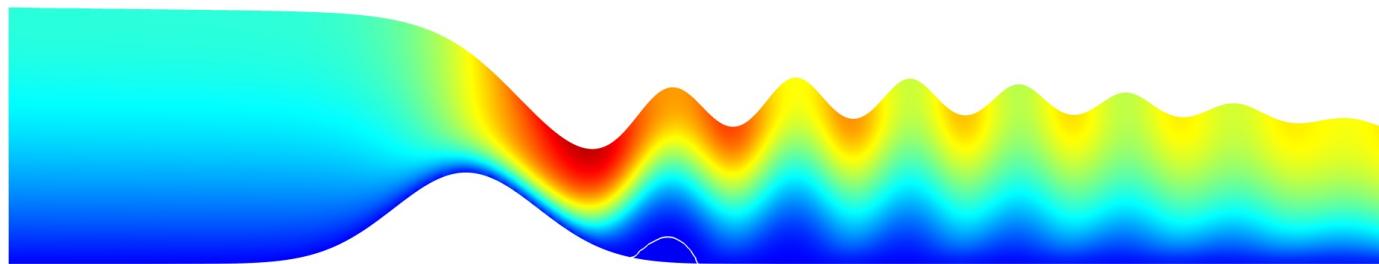


Navier–Stokes VOF

## Horizontal velocity

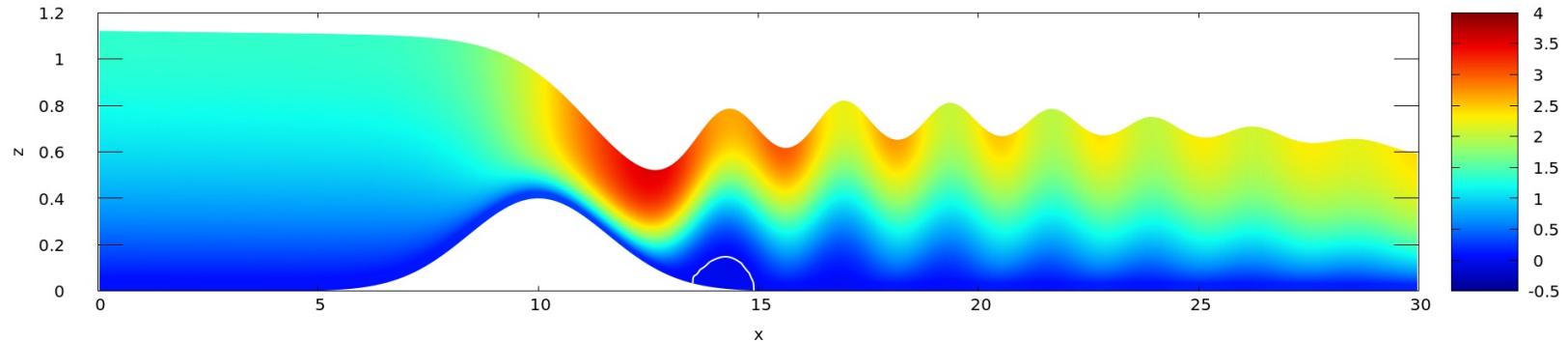


Multilayer non-hydrostatic (runtime: 3 min)

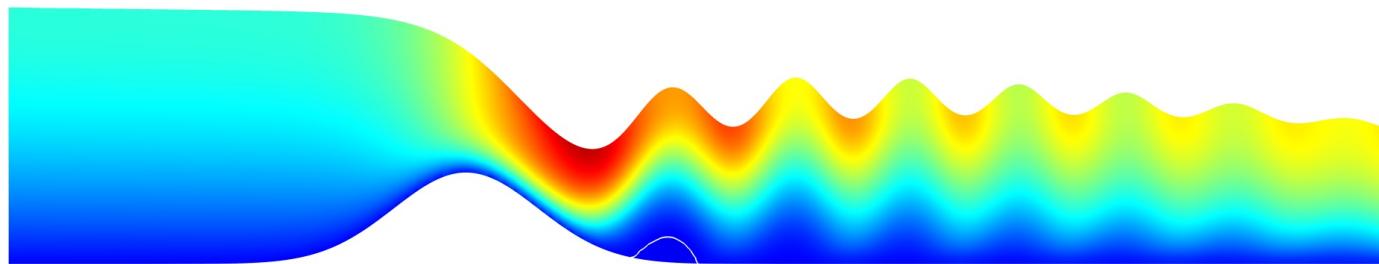


Navier–Stokes VOF

## Horizontal velocity



Multilayer non-hydrostatic (runtime: 3 min)

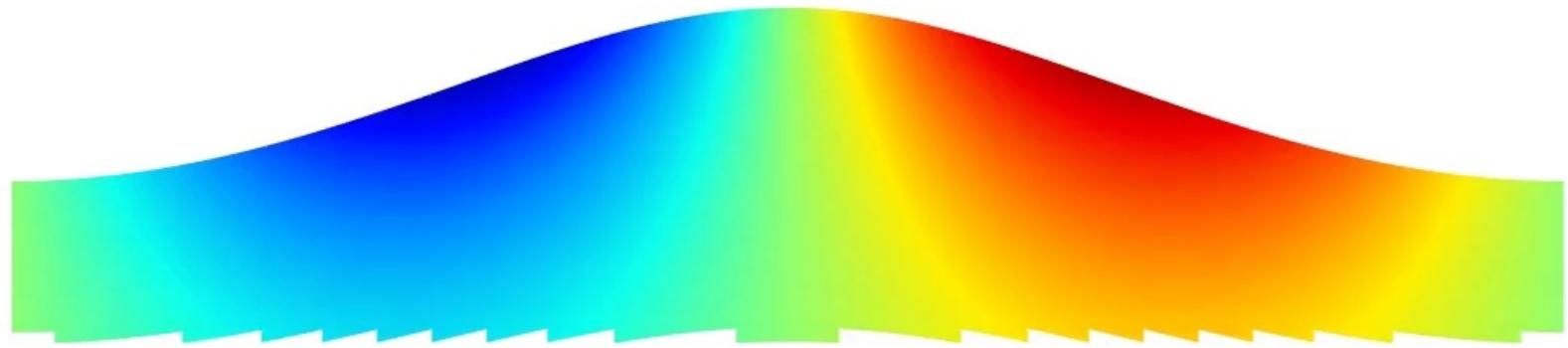
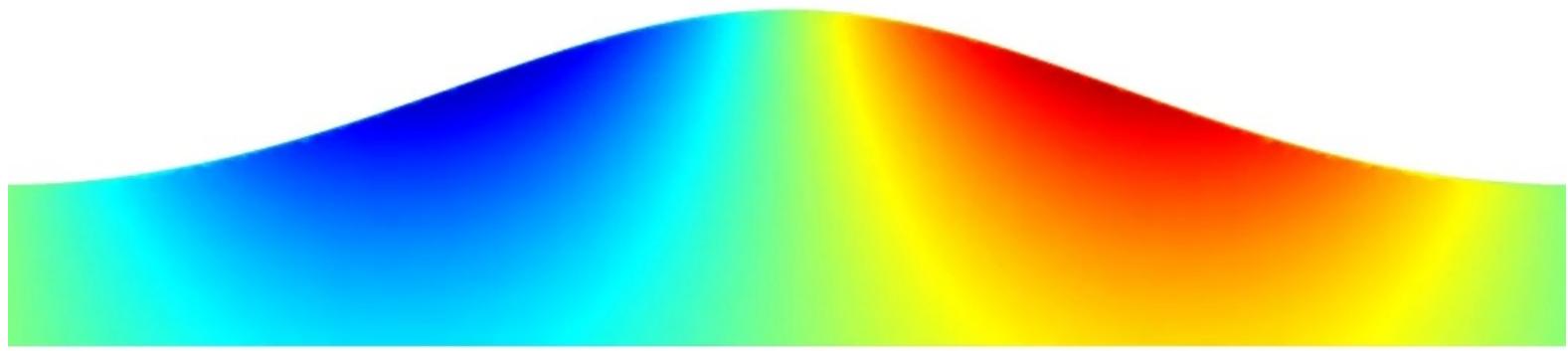


Navier–Stokes VOF (runtime: 1h15)

# Standing “river” waves in Waimea (Hawaii)

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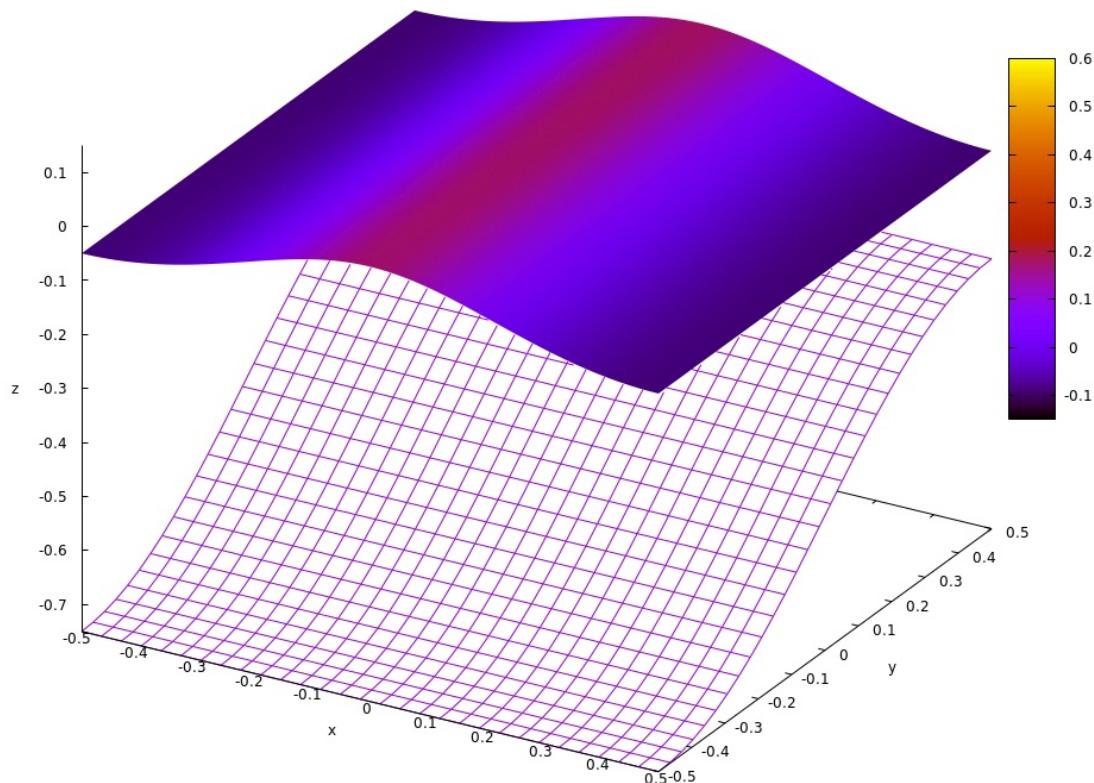


Vertical velocity: top: Navier–Stokes/VOF, bottom: multilayer

# 3D breaking Stokes wave

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$t = 0.00 \text{ TO}$



runtime: 13 minutes on 64 cores (MPI),  $256^2 \times 80$

- A semi-discrete consistent representation of the incompressible Euler/Navier–Stokes equations with a free-surface.
- This (new) set of equations has a clear physical interpretation and makes a seamless link between the Euler, Saint-Venant and Boussinesq equations.
- The same model gives accurate and efficient solutions for the evolution of metre-scale to kilometer-scale waves.
- Work in progress:
  - Multimaterial flows: densities, rheologies, surface tension etc.
  - Coriolis forces / geostrophic balance
  - Applications to ocean modelling
- Preprint submitted to JCP on HAL and [basilisk.fr](http://basilisk.fr)

The choice of a Lagrangian vertical coordinates implies a kinematic condition i.e.

$$w = \partial_t z + \mathbf{u} \cdot \nabla z$$

Note that this condition is physical only at the top and bottom boundaries. Using the vertical difference operator, the fact that  $h_k = [z]_k$  and the continuity equation

$$\nabla \cdot (h \mathbf{u})_k + [w - \mathbf{u} \cdot \nabla z]_k = 0$$

we get

$$\begin{aligned}[w]_k &= \partial_t [z]_k + [\mathbf{u} \cdot \nabla z]_k, \\ [w - \mathbf{u} \cdot \nabla z]_k &= \partial_t h_k, \\ -\nabla \cdot (h \mathbf{u})_k &= \partial_t h_k,\end{aligned}$$

which is the layer evolution equation.